

Variable density groundwater flow: are equivalent freshwater heads necessary or misleading?

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Variable Density Flow

A flow field containing fluids of different densities.



Velocity Potential [Muskat, 1937]

VS

Force Potential [Hubbert, 1940]

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Velocity potential

1. Energy related to unit volume
2. Underground fluids therefore are assumed to be incompressible
3. Equivalent freshwater heads are assumed to stand for the actual energy conditions in a flow field

Force potential

1. Energy related to unit mass

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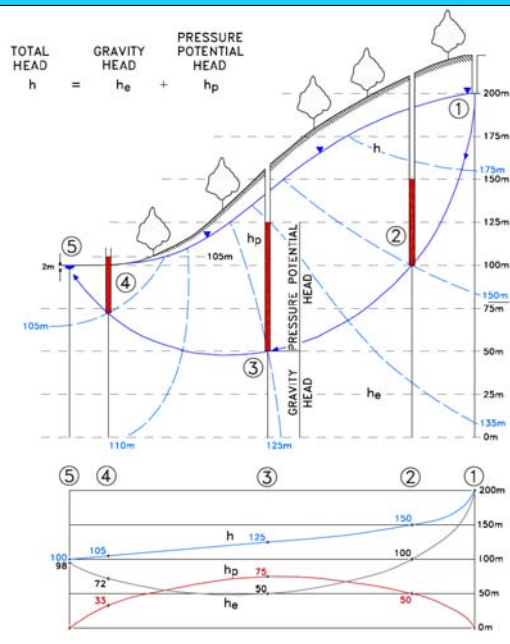


Gravitationally-Driven Groundwater Flow



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Unused energy stored by compressing the unit mass

Compression measured by water column in piezometers

after Weyer, 1978

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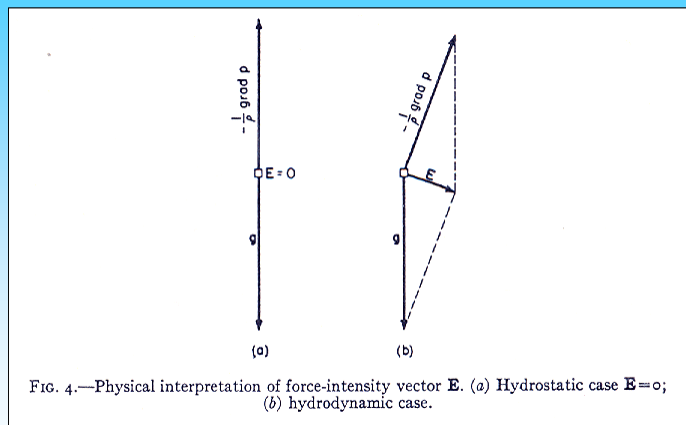
Buoyancy

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Hydrostatic vs Hydrodynamic

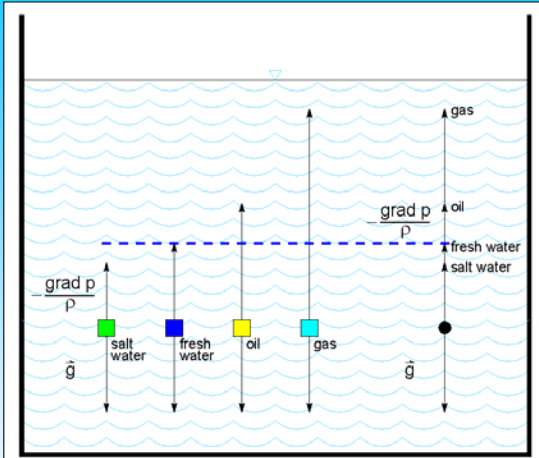


Hubbert, 1953. Entrapment of petroleum under hydrodynamic conditions

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- A = downward flow for dense salt water.... $(\text{grad } p)/\rho < \vec{g}$
- B = no flow of fresh water..... $(\text{grad } p)/\rho = \vec{g}$
- C = upward flow for oil..... $(\text{grad } p)/\rho > \vec{g}$
- D = upward flow for gas..... $(\text{grad } p)/\rho > \vec{g}$

after Weyer, 2010

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Schematic of hydrostatic forces on water, salt water, oil, and gas.

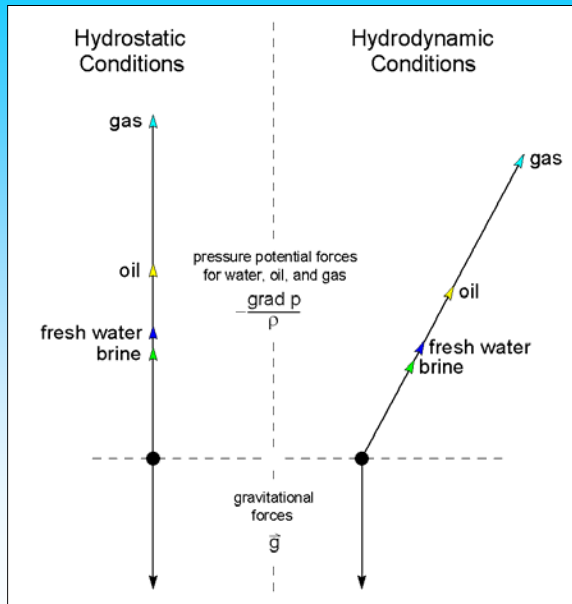
\vec{g} = gravitational force

$-(\text{grad } p)/\rho$ = pressure potential force (gradient of the pressure divided by the density)

buoyancy = $\vec{g} - (\text{grad } p)/\rho$

Hydrostatic Conditions

Hydrodynamic Conditions

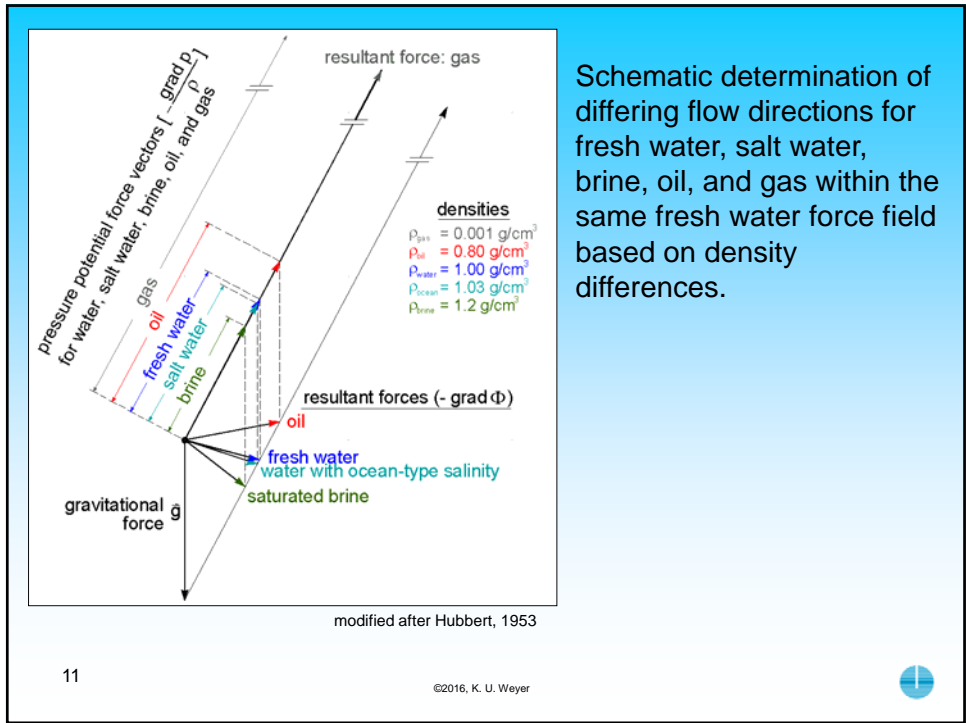


modified after Hubbert, 1953

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Schematic of hydrostatic vs hydrodynamic forces on water, oil, brine, and gas.



Schematic determination of differing flow directions for fresh water, salt water, brine, oil, and gas within the same fresh water force field based on density differences.

Field examples for variable density flow

Field example Salt River Basin, NWT, Canada:
Upward discharge to the surface of saturated brine.

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Upward migration of
saturated brine: Salt
River Basin, NWT,
Canada

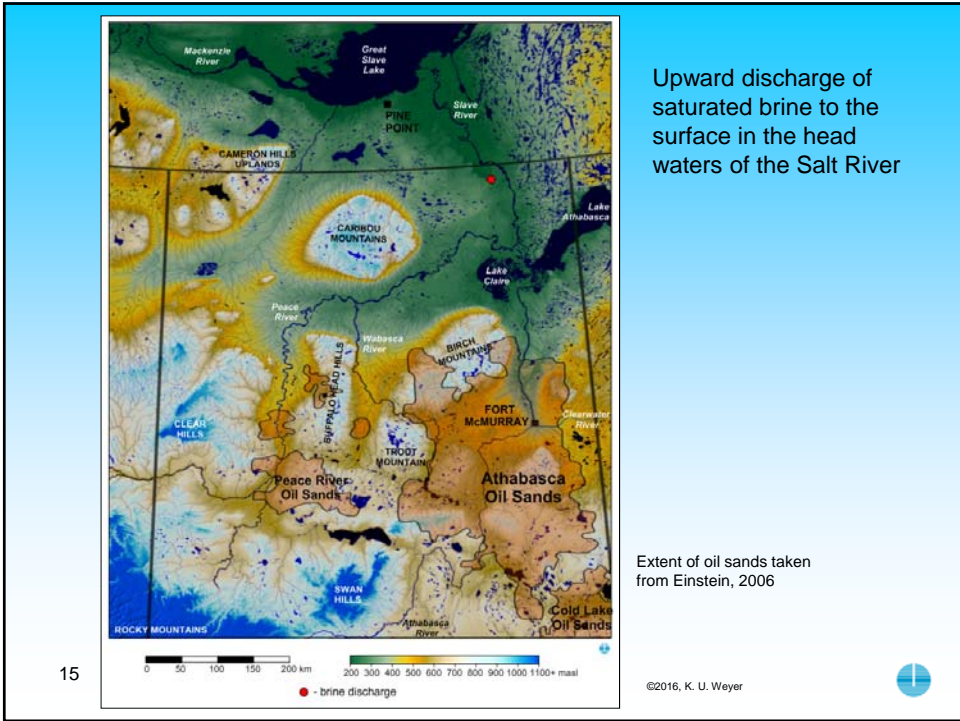
- Saturated [~ 350 g/l; density ~ 1.3 g/cm³] brine discharging upwards beside a creek
- Salt deposit is caused by precipitation of salt from brine not by evaporation of brine.

picture: K.U.Weyer, 1977

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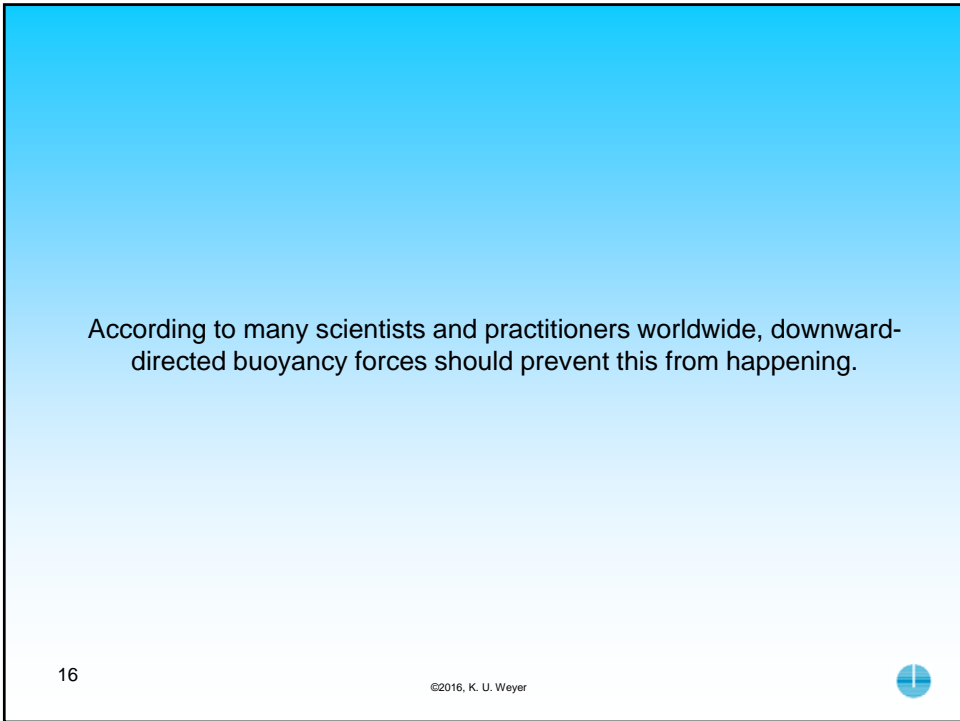
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Upward discharge of saturated brine to the surface in the head waters of the Salt River

According to many scientists and practitioners worldwide, downward-directed buoyancy forces should prevent this from happening.

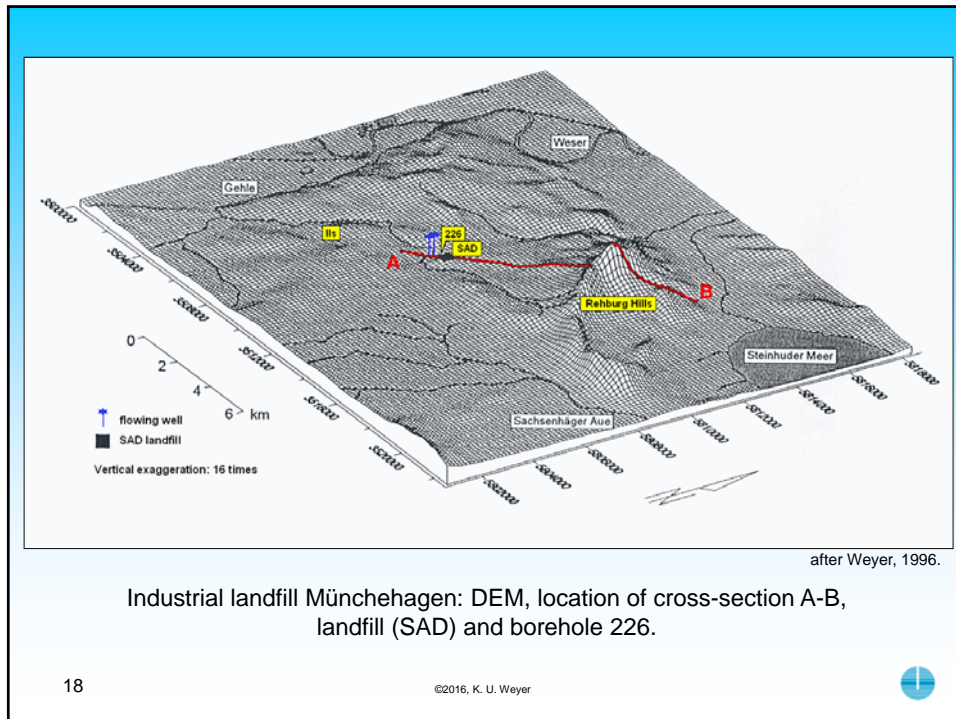


'Density changes' along flow lines within a real-world groundwater flow system

Field example and mathematical model: Upward discharge from depth of 1000 m of **ocean-type saline water** at Münchehagen, Germany

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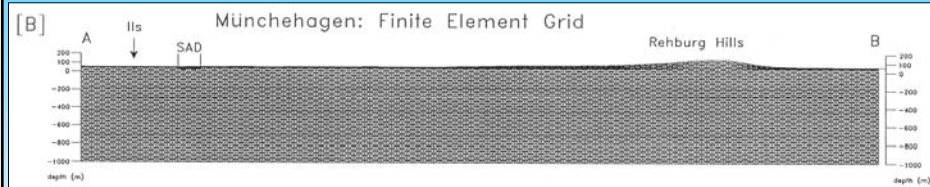


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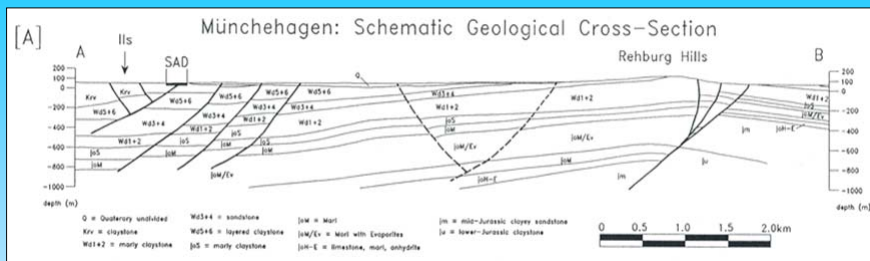


2D-vertical model of groundwater flow directions at the Münchehagen landfill area



Weyer and van Everdingen, 1995, Fig.3

Calculated groundwater flow directions based on groundwater table (following topography), geological structure, and estimated permeability contrasts. Program: FLONET



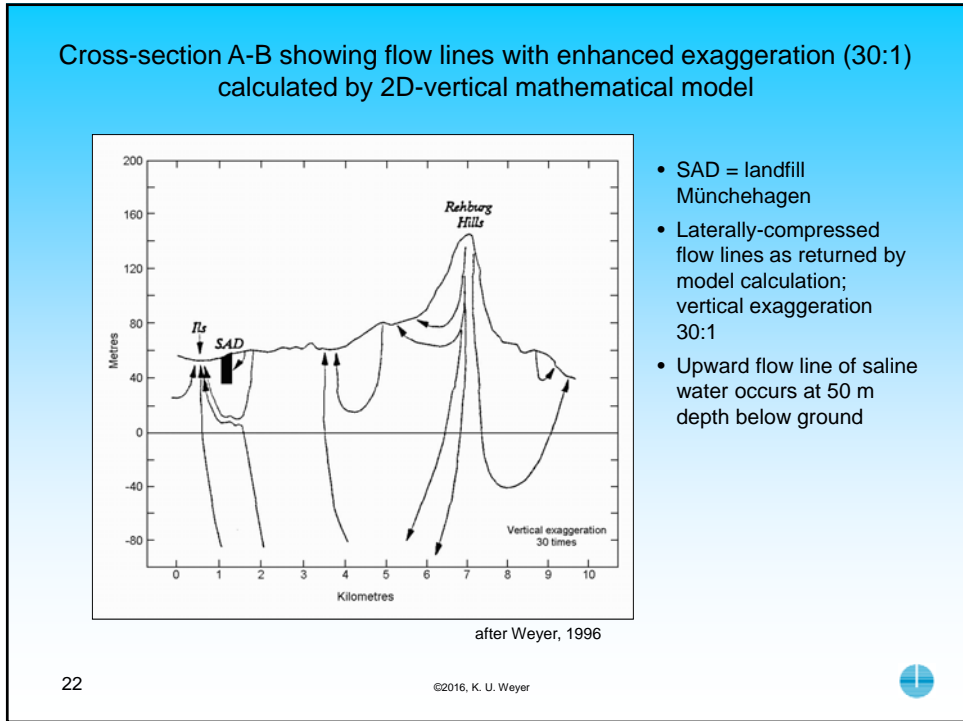
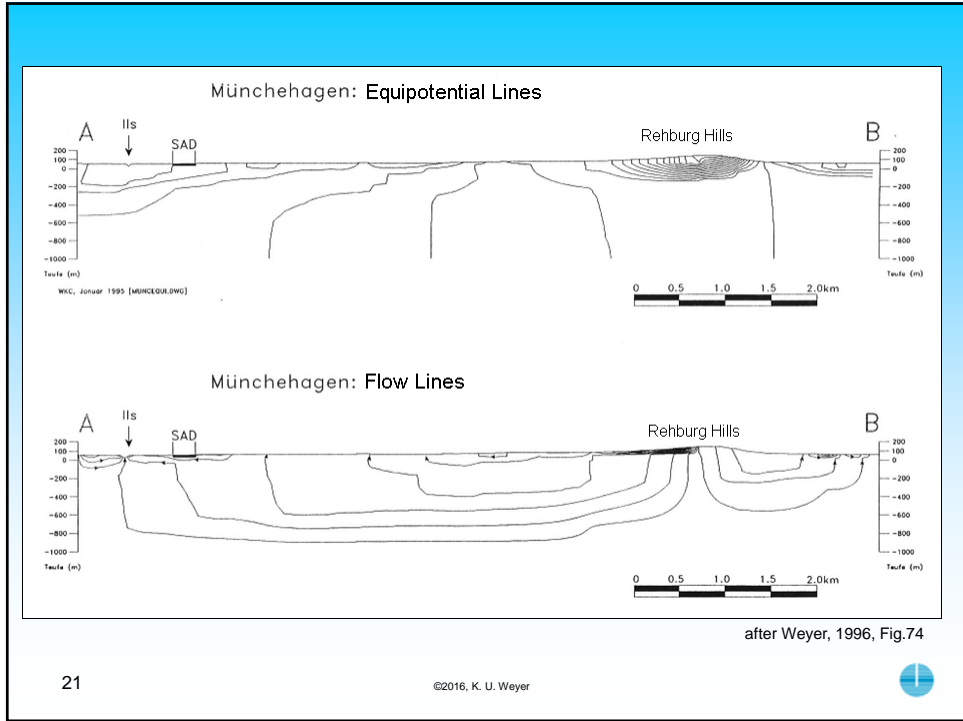
Weyer and van Everdingen, 1995, Fig.3 and Tab.1

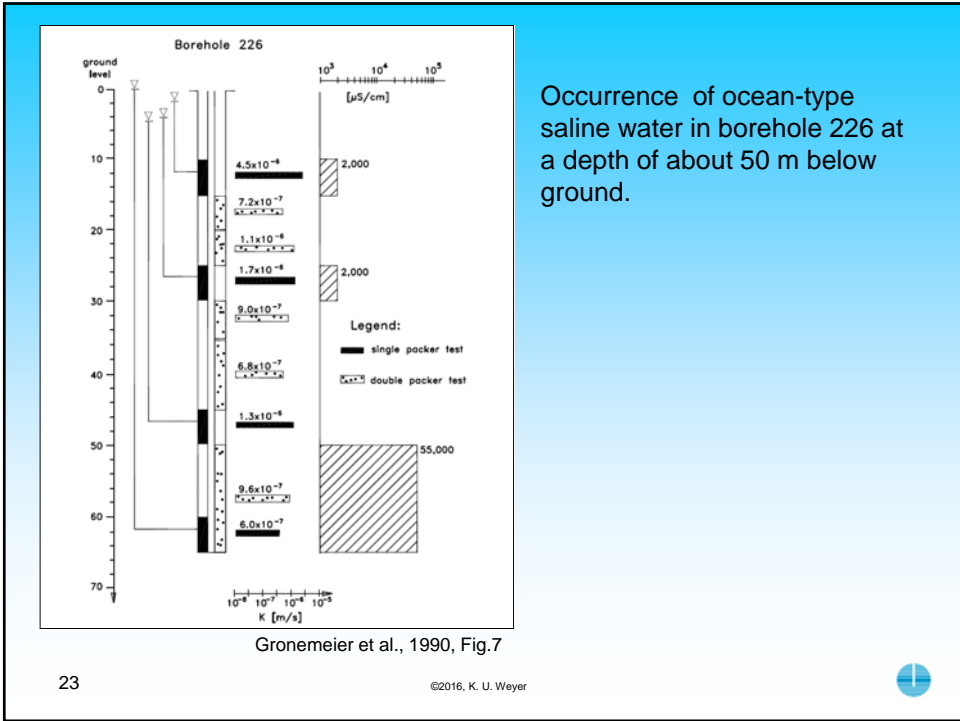
Table 1: Geology and assigned contrast permeabilities near Münchehagen

Formation Name	Abbr.	Geology	Permeability	
			Eastern half Loccum	Western half Rehburg
Quaternary	Q	sand, cobbles	—	—
Cretaceous	Kv	claystone 5-15m	200	—
	"	15-50m	100	—
	"	deeper Kv layers	50	—
Wealdon 1, 2	WD 1+2	marly claystone	5	1
Wealdon 3, 4	WD 3+4	sandstone	10	100
Wealdon 5, 6	WD 5+6	layered claystone	1	1
Serpult	joS	marly claystone	10	10
Mündener Mergel	joM	marl	100	100
Mündener Mergel	joM/Ev	marl with evaporites	100	100
Heersumer	joH-E	limestone, marl, anhydrite	100	100
mid-Jurassic	jm	clayey sandstone	10	10
low-Jurassic	ju	claystone	10	10

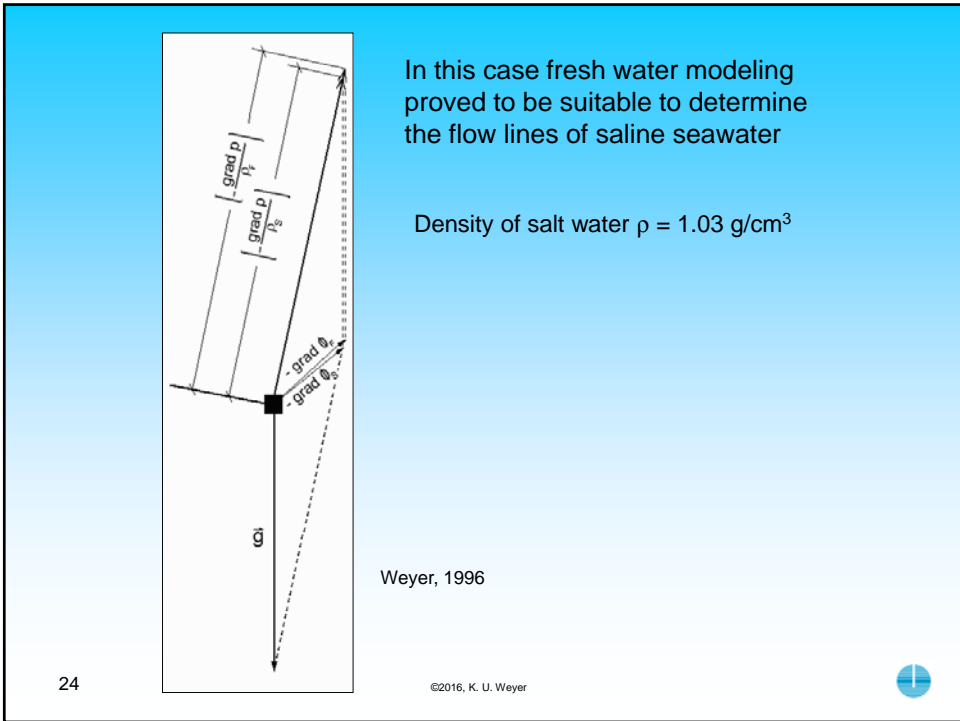
Geologic cross-section taken from official geologic maps 1:25,000 of the state of Lower Saxony, contrast permeabilities assigned by us.







Occurrence of ocean-type saline water in borehole 226 at a depth of about 50 m below ground.



Traditional ways of looking at variable density flow

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The big difference between assumptions and physical reality – or – where the handling of variable density flow usually goes astray

Referring to Bear (1972, p. 654), Bachu et al. (1993, p. 7) and Bachu and Underschultz (1993, p.1754) both considered two types of driving forces for variable density groundwater flow in northeast Alberta, “a ***potential force resulting from piezometric head differences***, and a ***buoyancy force resulting from density differences*** (Hubbert, 1940; Bear, 1972)”.

False quote: Hubbert (1940) does not refer to buoyancy forces or density differences

Bear, J., 1972. Dynamics of Fluids in Porous Media. American Elsevier Publishing Company, Inc., New York, NY, USA, 764p., ISBN: 978-0444001146

Bachu, S., J.R. Underschultz, B. Hitchon, and D. Cotterill, 1993. Regional-Scale Subsurface Hydrogeology in Northeast Alberta. Alberta Research Council, Bulletin No. 61, 44 p.

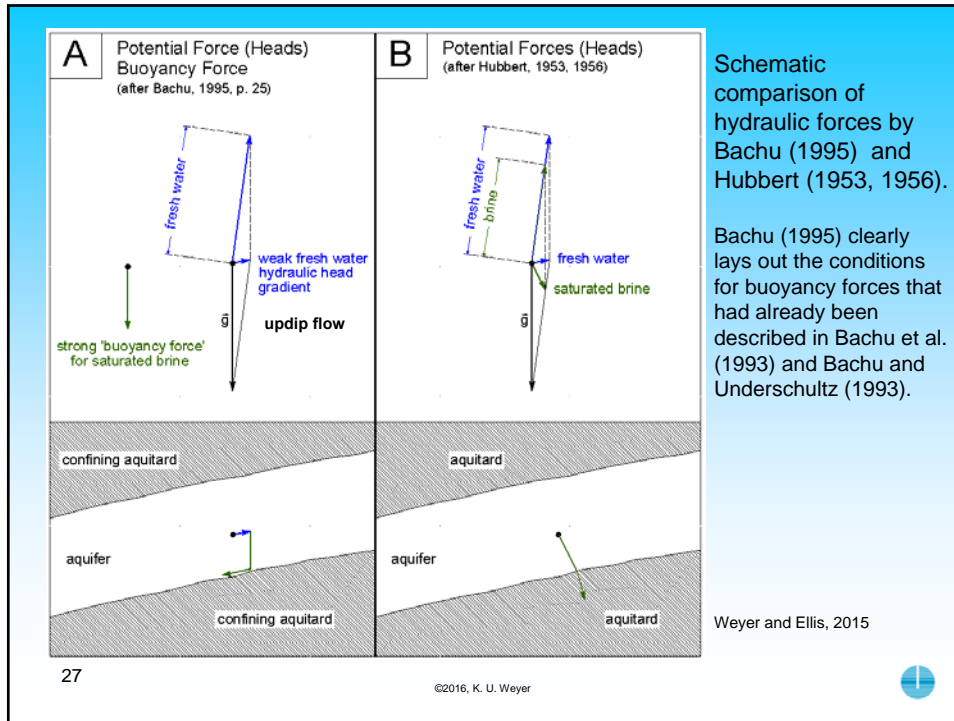
Bachu S., and J.R. Underschultz, 1993. Hydrogeology of formation waters, northeastern Alberta basin. AAPG Bulletin, vol. 77, issue 10, p. 1745-1768.

Bachu, S., 1995. Flow of variable-density formation water in deep sloping aquifers: review of methods of representation with case studies. Journal of Hydrology, vol. 164, p. 19-38.

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Schematic comparison of hydraulic forces by Bachu (1995) and Hubbert (1953, 1956).

Bachu (1995) clearly lays out the conditions for buoyancy forces that had already been described in Bachu et al. (1993) and Bachu and Underschlutz (1993).

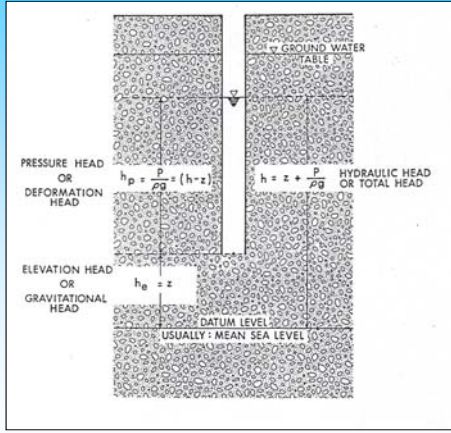
Equivalent freshwater head

A widely used attempt to determine the state of energy as a fresh water head within groundwater bodies of variable density

Heads

$$\Phi = h * g$$

potential energy = head • earth acceleration



- hydraulic head = elevation of water level in a piezometer (above datum level)
- gravitational head = screen elevation in gravitational field
- pressure head = height of water column in a piezometer

after Weyer, 1978



	Elevation z [m]	Pressure P [Pa]	Equiv. Fresh Water Head H _e = z + P/(rho _s *g) [m]	Gravitational Potential Phi _g = z*g [m ² /s ²]	Pressure Potential Phi _p = P/rho _s [m ² /s ²]	Total Force Potential (energy/mass) Phi = Phi _g + Phi _p [m ² /s ²]	Elevation Head H _e = Phi _g /g [m]	Pressure Head H _p = Phi _p /g [m]	Total Head H _t = Phi/g [m]
HYDROSTATIC	10	0	10	98.1	0	98.1	10	0	10
ocean-type salt water: rho _s = 1030 kg/m ³	9	10104.3	10.03	88.29	9.81	98.1	9	1	10
	8	20208.6	10.06	78.48	19.62	98.1	8	2	10
	7	30312.9	10.09	68.67	29.43	98.1	7	3	10
	6	40417.2	10.12	58.86	39.24	98.1	6	4	10
	5	50521.5	10.15	49.05	49.05	98.1	5	5	10
	4	60625.8	10.18	39.24	58.86	98.1	4	6	10
	3	70730.1	10.21	29.43	68.67	98.1	3	7	10
	2	80834.4	10.24	19.62	78.48	98.1	2	8	10
	1	90938.7	10.27	9.81	88.29	98.1	1	9	10
	0	101043	10.3	0	98.1	98.1	0	10	10

after Weyer and Ellis, 2015

Determination of equivalent freshwater heads in a tank filled with ocean type salt water. According to the equivalent freshwater head calculation the head at the bottom of the tank (10.3 m) is much higher than at the water surface (10 m) and upward flow must occur, which it does not. On the right of the table the correct head calculations are recorded according to force potential procedures confirming hydrostatic conditions with a head of 10 m throughout.



What we addressed:

- Hydrostatic vs hydrodynamic flow conditions (buoyancy forces).
- Bachu and Underschultz's (1993) buoyancy forces do not exist.
- In groundwater discharge areas, saturated brine can flow upwards to the surface.
- At the industrial landfill site Münchehagen variable density flow transports ocean-type salt water to the small river IIs. The respective flowlines were successfully calculated in a numerical model of a groundwater flow system using a density $\rho = 1 \text{ g/cm}^3$.
- Equivalent fresh water heads do not represent actual energy conditions in a variable density flow field and are misleading.

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Force Potential [energy/mass]

- Mass is measured in kilograms.
- A kilogram is independent of pressure, density, and temperature.
- Heads measured in piezometers of **any** pressure, density, temperature, etc., are the correct representation of energy if flow calculations and computer modelling are done with force potentials.
- Calculations and computer models with velocity potential lead to incorrect results in variable density flow.

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Synopsis

- When dealing with variable density flow under on-shore hydrodynamic conditions the application of physically-consistent force potentials and groundwater flow systems theory are the methods of choice.
- Computer programs making use of velocity potential (SUTRA and others) need to be surpassed by programs built on force potential. These programs do not yet exist.

